

Tides and Basins

Consider a basin with depth $H(\mathbf{x})$ forced by an equilibrium tide $\phi_e = g\eta_e$

$$\begin{aligned}\frac{\partial}{\partial t}\mathbf{u} + f\hat{\mathbf{z}} \times \mathbf{u} &= -\nabla(\phi - \phi_e) \\ \frac{\partial}{\partial t}\phi + \nabla \cdot [gH(\mathbf{x})\mathbf{u}] &= 0\end{aligned}$$

Let's solve for the velocities using

$$\begin{aligned}\mathbf{u}_t + f\hat{\mathbf{z}} \times \mathbf{u} &= -\nabla(\phi - \phi_e) \\ \text{(cross with) } \hat{\mathbf{z}} &\Rightarrow \\ \hat{\mathbf{z}} \times \mathbf{u}_t - f\mathbf{u} &= -\hat{\mathbf{z}} \times \nabla(\phi - \phi_e) \\ \text{(t derivative) } &\Rightarrow \\ \mathbf{u}_{tt} + f^2\mathbf{u} &= f\hat{\mathbf{z}} \times \nabla(\phi - \phi_e) - \frac{\partial}{\partial t}\nabla(\phi - \phi_e)\end{aligned}$$

Assuming $e^{-i\omega t}$,

$$(f^2 - \omega^2)\mathbf{u} = f\hat{\mathbf{z}} \times \nabla(\phi - \phi_e) + i\omega\nabla(\phi - \phi_e)$$

From $\hat{\mathbf{n}}$ dot this, we get the boundary conditions

$$-f\hat{\mathbf{t}} \cdot \nabla(\phi - \phi_e) + i\omega\hat{\mathbf{n}} \cdot \nabla(\phi - \phi_e) = 0$$

and the dynamical equation (multiplying the mass equation by $f^2 - \omega^2$)

$$\begin{aligned}i\omega(f^2 - \omega^2)\phi &= f\nabla \cdot [gH\hat{\mathbf{z}} \times \nabla(\phi - \phi_e)] + i\omega\nabla \cdot [gH\nabla(\phi - \phi_e)] \\ &= f\hat{\mathbf{z}} \cdot [\nabla(\phi - \phi_e) \times \nabla gH] + i\omega\nabla \cdot [gH\nabla(\phi - \phi_e)]\end{aligned}$$

so the general problem is

$$\begin{aligned}\nabla \cdot [gH\nabla(\phi - \phi_e)] &= (f^2 - \omega^2)\phi + \frac{if}{\omega}\hat{\mathbf{z}} \cdot [\nabla(\phi - \phi_e) \times \nabla gH] \\ f\hat{\mathbf{t}} \cdot \nabla(\phi - \phi_e) &= i\omega\hat{\mathbf{n}} \cdot \nabla(\phi - \phi_e)\end{aligned}$$

If we work in a circular basin with $H = H(r)$ and take an $e^{im\theta}$ form, this simplifies to

$$\begin{aligned}\nabla \cdot gH\nabla(\phi - \phi_e) &= (f^2 - \omega^2)\phi + \frac{fm}{\omega a}gH_r(\phi - \phi_e) \\ \frac{\partial}{\partial r}(\phi - \phi_e) &= \frac{fm}{\omega a}(\phi - \phi_e)\end{aligned}$$

Numerical note: we would probably want to solve in terms of the height anomaly $\eta = \phi - \phi_e$ which has homogeneous but mixed boundary conditions and satisfies a forced equation in the interior.

Constant depth

When gH is constant, we have

$$\begin{aligned}\nabla^2(\phi - \phi_e) &= \frac{f^2 - \omega^2}{gH}\phi \\ \frac{\partial}{\partial r}(\phi - \phi_e) &= \frac{fm}{\omega a}(\phi - \phi_e)\end{aligned}$$

We can write the equations in non-dimensional form, scaling distances by the domain scale a , and ω by f . Then

$$\begin{aligned}\nabla^2(\phi - \phi_e) &= \alpha^2(1 - \omega^2)\phi \\ \frac{\partial}{\partial r}(\phi - \phi_e) &= \frac{m}{\omega}(\phi - \phi_e)\end{aligned}$$

with $\alpha = a/R_d$ and $R_d = \sqrt{gH}/f$.

FREE MODES: when $\phi_e = 0$, we have an eigenvalue problem for ω . Let us make the *ansatz* that $\omega > 1$ (frequency greater than f – super-inertial). Then

$$\phi = J_m(Kr) \quad , \quad K^2 = \alpha^2(\omega^2 - 1)$$

This gives one equation relating K to ω ; the other is

$$K J'_m(K) = \frac{m}{\omega a} J_m(K) \quad \text{or} \quad \frac{K J'_m(K)}{J_m(K)} = \frac{m}{\omega}$$

For $\omega \gg 1$, the boundary condition says the K is a zero of J'_m ; the “free-end” version of the drumhead modes (i.e., $\phi_r \simeq 0$ rather than $\phi = 0$). We expect this to happen for small basins since the time for a gravity wave to cross a/\sqrt{gH} becomes short compared to $1/f$. The dispersion relation

$$\omega = \pm \sqrt{1 + (K/\alpha)^2}$$

has a large parameter multiplying K (which itself is greater than 2.4)

For the low frequency case,

$$\phi = I_m(Kr) \quad , \quad \omega^2 = 1 - \frac{K^2}{\alpha^2} \quad \text{and} \quad \frac{K I'_m(K)}{I_m(K)} = \frac{m}{\omega}$$

For the previous case, ω could have both positive and negative values since the J_m functions are oscillatory. Now, however, I_m is strictly increasing so that the left-hand side is positive and ω will be positive. The waves, will propagate counter-clockwise around the basin.

They are Kelvin waves with the boundary on their right in the northern hemisphere. If ω is small enough, K will be large and $I'_m/I_M \rightarrow 1$; i.e.,

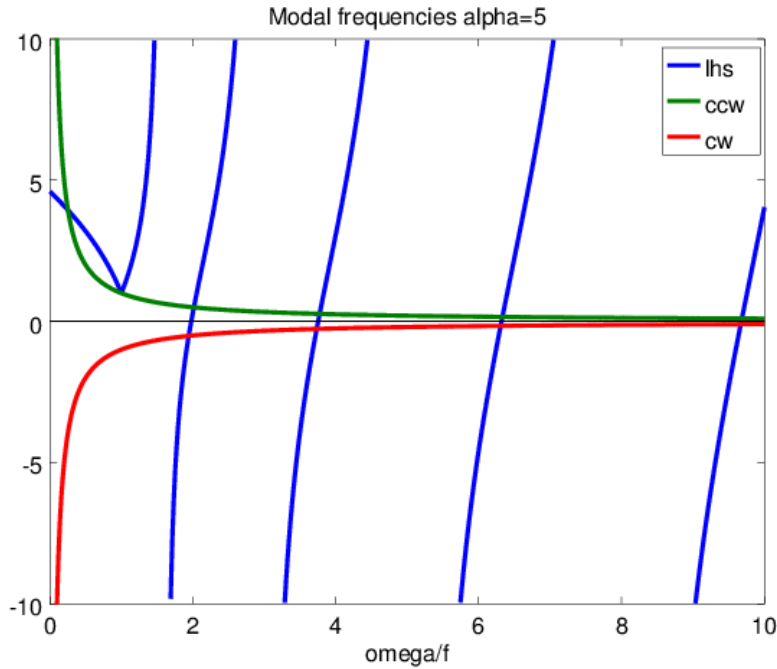
$$\omega \simeq m/K \simeq m/\alpha$$

In dimensional form, this is

$$\omega = fmR_d/a = (m/a)\sqrt{gH}$$

making the Kelvin wave characteristic clear, since m/a is the wavenumber.

We show the left and right sides of the boundary condition equations for $m = 1$;



the frequencies are given by the intersections with either the clockwise ($w < 0$) or counterclockwise ($w > 0$) curves. If $\omega < f$, we have just the single ccw Kelvin mode; above f , the two Poincaré waves have quite similar frequencies. The successive roots correspond to more and more zeros in the radial direction.

There is also appears to be an inertial wave $\omega = f$; in this limit $K \rightarrow 0$ and the solution to the interior equation is just $\phi = r^m$; this is consistent with the boundary condition $\phi_r(a) = m\phi(a)/a$. However, you cannot generally find a consistent radial flow structure that satisfies the boundary conditions; this is a spurious mode.

FORCED MODES: in the simple case where the basin is small enough, we can take $\phi_e = \phi_0 x \cos(\omega t) = \phi_0 r \cos(\theta) \cos(\omega t)$; this has $\nabla^2 \phi_e = 0$ so that the interior equation is still

$$\nabla^2 \phi = \alpha^2 (1 - \omega^2) \phi$$

However ω is set by the tidal forcing, as is $m = 1$, so we will have I_1 solutions for sub-inertial frequencies and J_1 for super-inertial cases. The boundary conditions now determine the amplitude of the waves. Since

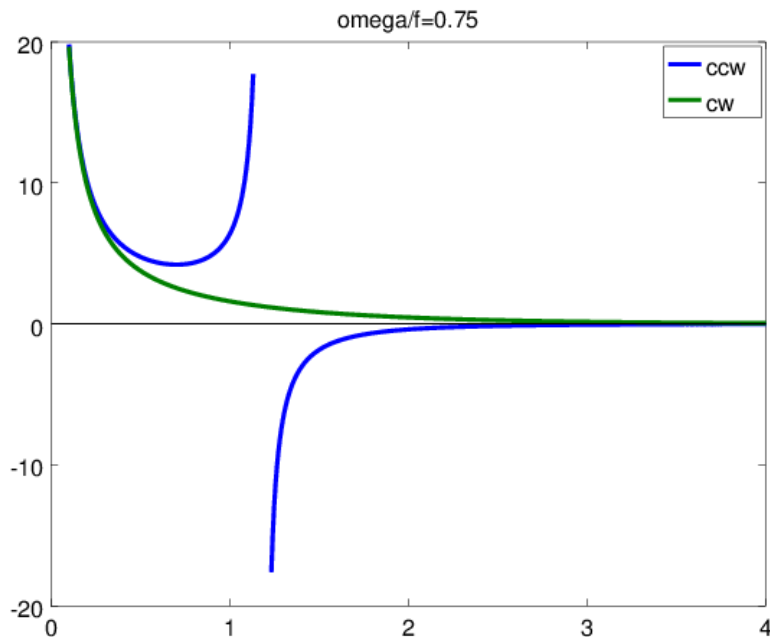
$$\cos(\theta) \cos(\omega t) = \frac{1}{2} [\cos(\theta + \omega t) + \cos(\theta - \omega t)]$$

we are forcing waves in both directions around the boundary. In terms of $\Re \exp(i\theta - i\omega t)$ we have both positive and negative frequencies, forced with equal amplitudes.

For subinertial cases, $\phi = A \frac{\phi_0}{2} I_1(Kr)$, we find

$$A [I_1(K) \frac{1}{\omega} - K I_1'(K)] = \frac{1}{\omega} - 1$$

Here $K = \alpha \sqrt{1 - \omega^2}$.



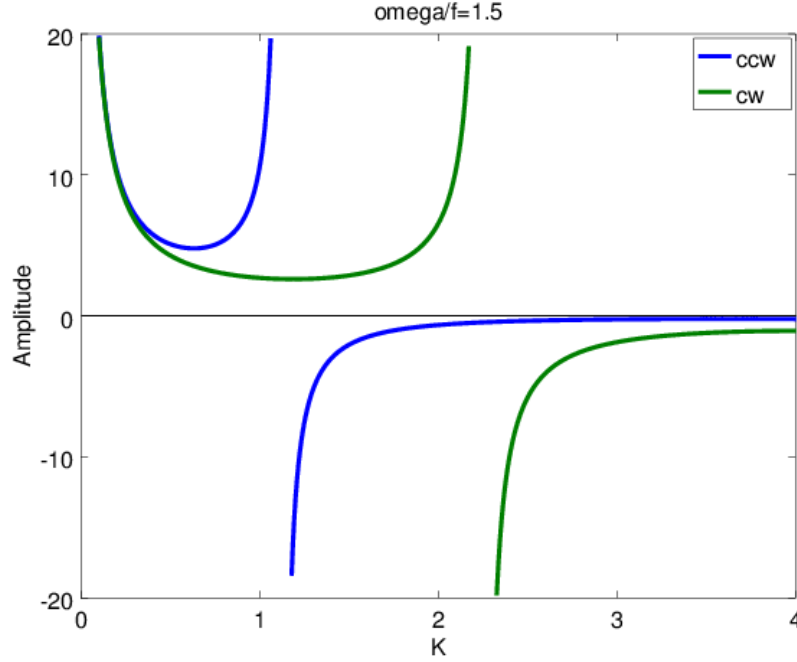
Amplitude as a function of basin size for subinertial motions $\omega = 0.75$

For the subinertial case, with I_1 solutions, the (single) singularity only occurs for the ccw ($n = 1$) solution since I_1' and I_1 are both positive.

The superinertial case has roots at all the free modes, both cw and ccw since J_1 and J_1' take on both signs..

$$A[J_1(K)\frac{1}{\omega} - KJ_1'(K)] = \frac{1}{\omega} - 1$$

with $K = \alpha\sqrt{\omega^2 - 1}$.



Amplitude as a function of basin size for superinertial motions $\omega/f = 1.5$

Note that the modes here are widely separated because the plots are for a small basin $\alpha = 1$.

Bowl

When $H = H(r)$, eqn (eq) is an ODE in r and should be solvable numerically. Free solutions should now include two gravity waves and a topographic Rossby wave; these should propagate ccw around the basin. So let's check that: suppose $H = H_0[1 - sr^2/2a^2] = H_0h$ so that $H'/H_0 = -sr/a^2$, then

$$\nabla \cdot h\nabla\phi = \frac{f^2 - \omega^2}{gH_0}\phi - \frac{fns}{\omega a^2}\phi$$

($h = 1 - s^2r^2/2a^2$) with

$$\omega \frac{\partial}{\partial r}\phi = \frac{fn}{a}\phi$$

at $r = a$. We approximate $\omega \ll f$ and $s \ll 1$ to get

$$\nabla^2\phi = \frac{f^2}{gH_0}\phi - \frac{fns}{\omega a^2}\phi$$

with $\phi(a, \theta) = 0$. This has solutions

$$\phi = J_n(k_n r/a)$$

with k_n a zero of the Bessel function and

$$-k_n^2 - \frac{f^2}{gH_0} = -\frac{fn s}{\omega a^2}$$

for $\omega > 0$ we have the $n > 0$ mode (ccw) with the shallow water on the right. Note that, for the forced problem, the solutions for $n = \pm 1$ will no longer be the same.